

**Lithostratigraphy and Diagenesis of the Upper Pennsylvanian (Missourian)
Lansing-Kansas City Groups in Rooks County, Kansas**

By

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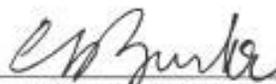
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Abstract

The Upper Pennsylvanian (uppermost Missourian) Lansing-Kansas City Groups (LKC) are prime oil reservoirs targeted by many petroleum companies exploring in Kansas and Nebraska. The LKC groups are a series of cyclothems with carbonate and terrigenous clastic components. A 240 foot, full-diameter core from within Dopita Field in Rooks County, Kansas that recovered the entire LKC section, along with thin sections and plugs (porosity and permeability tested), were studied to determine the depositional facies and diagenetic controls on reservoir development in these rocks, and to understand which producing zones in the study area, if any, would be candidates for a waterflood. The LKC section in the core (zones labeled A-L from top to bottom) was divided into lowstand-transgressive-highstand complete cycles and incomplete sub-cycles. A complete cycle consists of, from base to top: (i) a section with evidence of subaerial exposure, (ii) a transgressive-systems tract and ensuing maximum flooding surface, (iii) a regressive highstand-systems tract, and (iv) another section suggestive of subaerial exposure. An incomplete sub-cycle is represented by the upward change from transgression to highstand without evidence of subaerial exposure. Throughout the core there were 6 complete cycles and 10 incomplete sub-cycles identified. The study showed that any porosity and permeability (along with oil shows) present in the rocks had developed during times of subaerial exposure within shallow-water carbonate sands deposited in the highstand systems tracts of cycles. Currently plans are being made to begin waterflooding the south end of the Dopita Field. The results from this study are being implemented in deciding which zones to flood.

INTRODUCTION

The Lansing-Kansas City Groups comprise stacked cyclothems with locally prolific hydrocarbon-bearing carbonate reservoirs. They are sought-after throughout the midcontinent and are especially prolific on the Central Kansas Uplift and the Cambridge Arch (Figure 1, Figure 2). The Lansing-Kansas City groups are known not only to produce on structural highs but also areas off-structure where individual zones within the section thicken locally (Watney, 1980). Many regional studies of these rocks in the midcontinent have been done, primarily in Kansas and Nebraska (e.g., Parkhurst, 1959, 1962; Rascoe and Adler, 1971; Heckel and Baeseman, 1975; Heckel, 1980, 1984, 1986, 1995; Brown, 1984; Watney et al., 1989; Heckel et

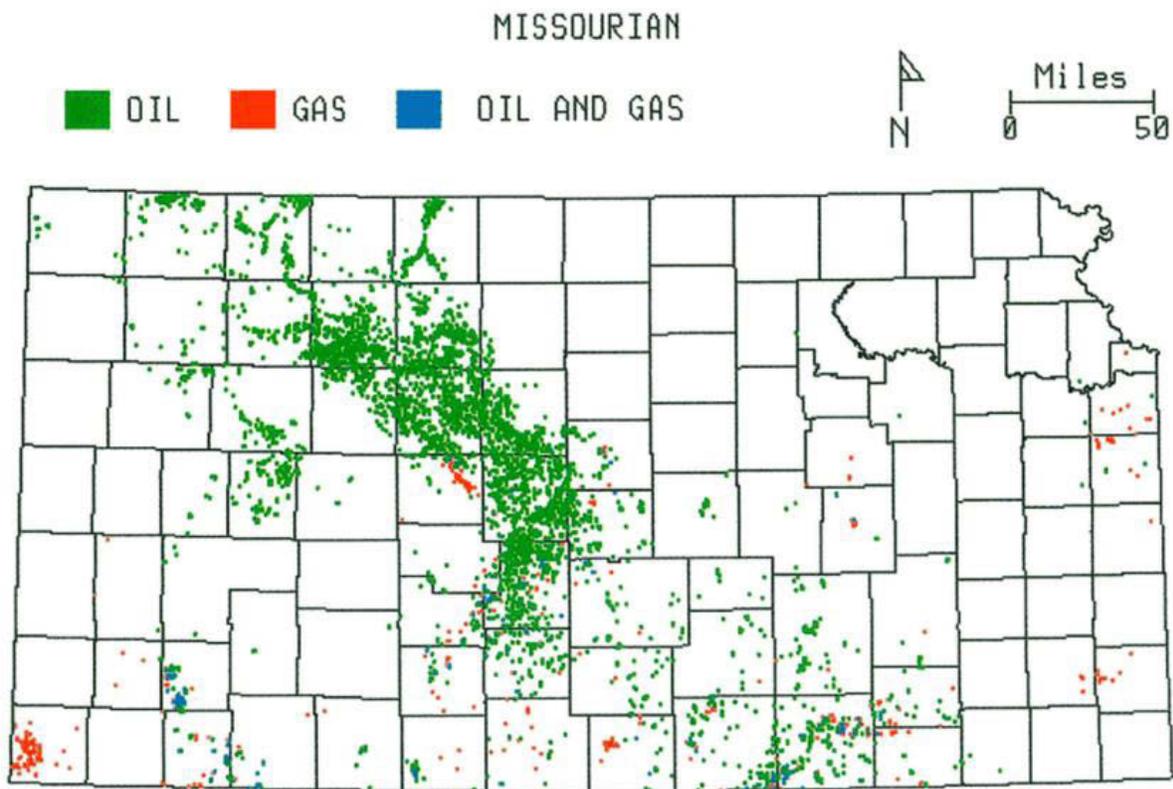


Figure 1 – Distribution of petroleum production in the Lansing-Kansas City/Pleasanton groups in Kansas (KGS, 1987).

al., 2007;).

Group	Stage	Series
WABAUNSEE	VIRGILIAN	UPPER PENNSYLVANIAN
SHAWNEE		
DOUGLAS		
LANSING	MISSOURIAN	
KANSAS CITY		
PLEASANTON		
MARMATON	DESMOINESIAN	MIDDLE PENN.
CHEROKEE		

Figure 2 - Stratigraphy of the Lansing-Kansas City groups (from Zeller, 1968).

The purpose of this study is to get a better understanding of the lithological and diagenetic controls on reservoir development and continuity in the Lansing-Kansas City section in Dopita Field, with the end result aiding in a secondary recovery by waterflood.

Study Area

The study area, and the location of the core examined, is in the southern half of the Dopita Field on the northeastern edge of the Central Kansas Uplift, in approximately the center

of Rooks County, Kansas. The field is about 10 sq. miles in size with approximately 180 wells drilled. Cumulative production of 13.8 million barrels is primarily out of the Arbuckle Group, and many wells in the field were later recompleted in, and new wells drilled into, the Lansing-Kansas City groups (Oil, 2013).

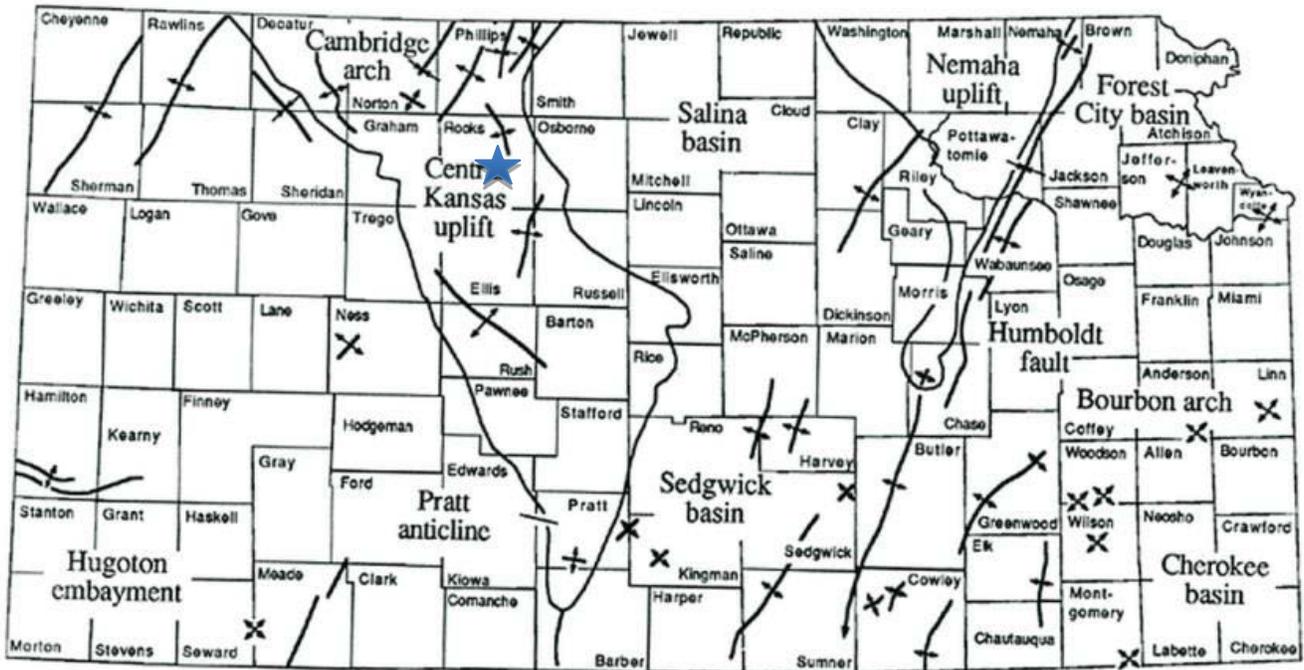


Figure 3 – Post-Mississippian structure map of Kansas (from Merriam, 1963). Blue star represents the location of the core used for this study.

General Geology of the Area

The study area is on the northern to northeastern flank of the Central Kansas Uplift, just southeast of the Cambridge arch and east of the Salina Basin (Figure 2). Both of these

features came into existence during the pre-Morrowan Epeirogeny (pre-Pennsylvanian) (Rascoe and Adler, 1983). With structural folding and probable sea-level drop, Mississippian rocks were eroded off the Cambridge Arch and most of the Central Kansas uplift (Rascoe and Adler, 1983). In early Desmoinesian (uppermost Middle Pennsylvanian) time, the Arch and Uplift were clastic hills and were transgressed by seas through late Desmoinesian time (Rascoe and Adler, 1983). During Missourian time the Central Kansas Uplift and Cambridge Arch were transgressed by epeiric seas several times, likely caused by glacial eustasy (Boardman and Heckel, 1989). The Lansing-Kansas City groups on the Central Kansas Uplift were deposited during the Late Pennsylvanian (Missourian), a time of numerous eustatic sea level fluctuations (Boardman and Heckel, 1989). These eustatic fluctuations and resulting transgression and regression were likely the result of glacial buildups and alternating meltings (e.g., Wanless and Shepard, 1936; Crowell, 1978; Boardman and Heckel, 1989). Heckel (1980, 1984, 1994) ruled out the alternative explanations of local tectonism or migrating deltas that had been used previously to explain these cycles. Veevers and Powell (1987) found evidence that the glacial deposits on Gondwanaland (the megacontinent to the south of the study area) reached a peak during the Middle to Late Pennsylvanian time, and such glaciations and deglaciations presumably affected eustasy in the study area (Veevers and Powell, 1987). The periodicity and frequency of the cyclothems in the middle to late Pennsylvanian generally are ascribed to Milankovitch cycles such as those suggested for during the Pleistocene -- that is, one cycle per 20,000 to 40,000 years up to 100,000 to 400,000 years (e.g., Veevers and Powell, 1987).

Reservoir and Log Characteristics

The LKC zones in each well in the southern end of Dopita Field were correlated and mapped using gamma ray-density/neutron logs. A structure map was prepared for the top of the LKC and the base of the KC, an isopach map was made for the entire LKC, and Porosity*Feet “pay zone” maps also were prepared for every zone in the LKC. These maps were prepared prior to coring so that an optimum core location could be decided on prior to drilling and obtaining the core. Percent porosity and zone thickness were used to estimate the Porosity*Feet values for the well logs. Only clean limestones (according to the gamma ray logs) with porosity of at least 5% were given values. The thickness in feet of the clean limestone in every LKC zone was multiplied by the average porosity through each individual zone to get an effective value of Porosity*Feet in the units mapped. Despite these maps, the cored well ran significantly lower and porosity values also were lower with respect to surrounding wells.

Methods of Study

Besides basic log analysis, a full-diameter core containing the entire LKC groups (240 feet), 7 representative thin sections, and 16 permeability plugs (tested for porosity and permeability) were analyzed from Trek AEC LLC’s Dopita A #16 well in Rooks County, Kansas. As noted above, the well ran low to surrounding wells, and after a failure to produce the LKC ‘G’ and ‘I’ zones, the current plan is to make the well an injection well as part of the secondary recovery program that is in the process of being implemented.

The core was extracted by Devilbiss Coring Service, Inc. on W-W Drilling Rig #6. The core was taken in one-stand increments, with two 30 ft. core barrels recovering 60 ft. at a

time. There were 4 total stands taken, with full recoveries on all 4 stands, resulting in 240 ft. of continuous core; the entire LKC groups and partial overlying and underlying units were recovered. The core was kept in aluminum sleeves cut in 3 ft. sections, capped, and transported back to Wichita State University Geology Department's rock preparation room in Wichita, Kansas. The core was personally slabbed by this author to minimize the potential for saw blade marks often left behind by major labs, then it was boxed and placed in the Carbonate Laboratory at WSU. Here, the slab faces were polished and then acidized (with 10% HCl) for a clean face with high relief. The prepared slabs were then analyzed using a binocular microscope. Alizarin Red S stain was used to distinguish between limestones and dolomites, and more dilute acid was used as needed. Dunham's (1962) classification of carbonate rocks was used to describe the lithologies of the core.

Seven representative thin sections were taken from two zones with production potential -- the LKC 'G' and 'I' zones. The thin sections were made using a blue-dyed epoxy to make any porosity readily visible, and they were examined using a standard petrographic microscope.

Sixteen 1" permeability plugs, each approximately 3" long, were drilled with a hole saw from the butt ends of the core. The depths selected for permeability plugs were based on the presence of porosity, oil staining, and degree of oil saturation. These plugs were sent to Weatherford Laboratories in Houston, Texas for a routine porosity and permeability analyses following standard analytical procedures. After extracting all the fluids, the plugs were vacuum dried at 180°F and tested with a net confining stress of 800 psi.

Lithostratigraphy

Cycles within the LKC section elsewhere in Kansas have been described by Watney et al. (1980) and others (Prather, 1984) as having 4 basic components: (i) a thin but distinctive basal transgressive unit deposited as sea level rose, (ii) overlain by a marine shale, (iii) overlain by a regressive carbonate and, in turn, (iv) a regressive shale deposited during sea level fall. The same cycle components were identified in the Dopita Field core, although differences in lithology and complexity of the cycles in the core were further evaluated to determine if any finer sedimentological details that may be present had relevance to petroleum reservoir development in the section. For example, some cycles in the core would indicate shallowing to the point of subaerial exposure, while other cycles included only a shallow-marine shale or tidal-flat deposit at the tops of (incomplete) cycles. For this study, the widely industry-accepted 'A' through 'L' zonal identification scheme of the LKC section was initially ignored and a sea-level curve was developed based on identification of cycles in the core. Afterwards, the zones within the LKC section were added to these interpretations (Boardman and Heckel, 1989).

When developing a columnar lithologic section for the core an attempt was made to distinguish more complete cycles from relatively more minor, incomplete cycles to determine what role cycle completeness might exert on petroleum reservoir (porosity) development. It was determined in the present study that a complete cycle in the core typically includes transgressive marine deposits overlying rocks affected by subaerial exposure (such as paleosols), and such transgression generally culminated in deposition of inferred maximum flooding surface deposits (generally black, phosphatic shale or dark gray shale)(Boardman et al., 1984). These deposits comprise the transgressive-systems tracts of LKC cycles. They are in turn overlain by regressive, upward-shallowing carbonate rocks that are micritic at the base and which pass upward into inferred high-energy biograinstone to packstone or oolite packstone to grainstone (carbonate

sands) that comprise the highstand-systems tracts of LKC cycles. These highstand-systems tract deposits generally are overlain by paleosols in red mudrocks indicative of subaerial exposure (Figure 4-A). Incomplete, relatively minor sub-cycles include similar transgressive-systems tracts limestones overlain by upward-shallowing limestones, locally with some tidal-flat deposits (Figure 4-B-C), and without capping paleosols or only very poorly developed (incipient) paleosols in green mudrocks (Figure 4-D). Similar tidal-flat deposits and incipient paleosols have been described by many workers, including for example Shinn (1983) and Retallack (1988), respectively. Soilstone crusts were interpreted as inferred slight unconformities (Figure 5), and sediment accumulation beneath grains was interpreted as being caused by an abrupt sea level drop (Figure 6)(Shinn, 1983). The high-energy packstones and grainstones make up the reservoir rock for all the zones (Figures 7-12).

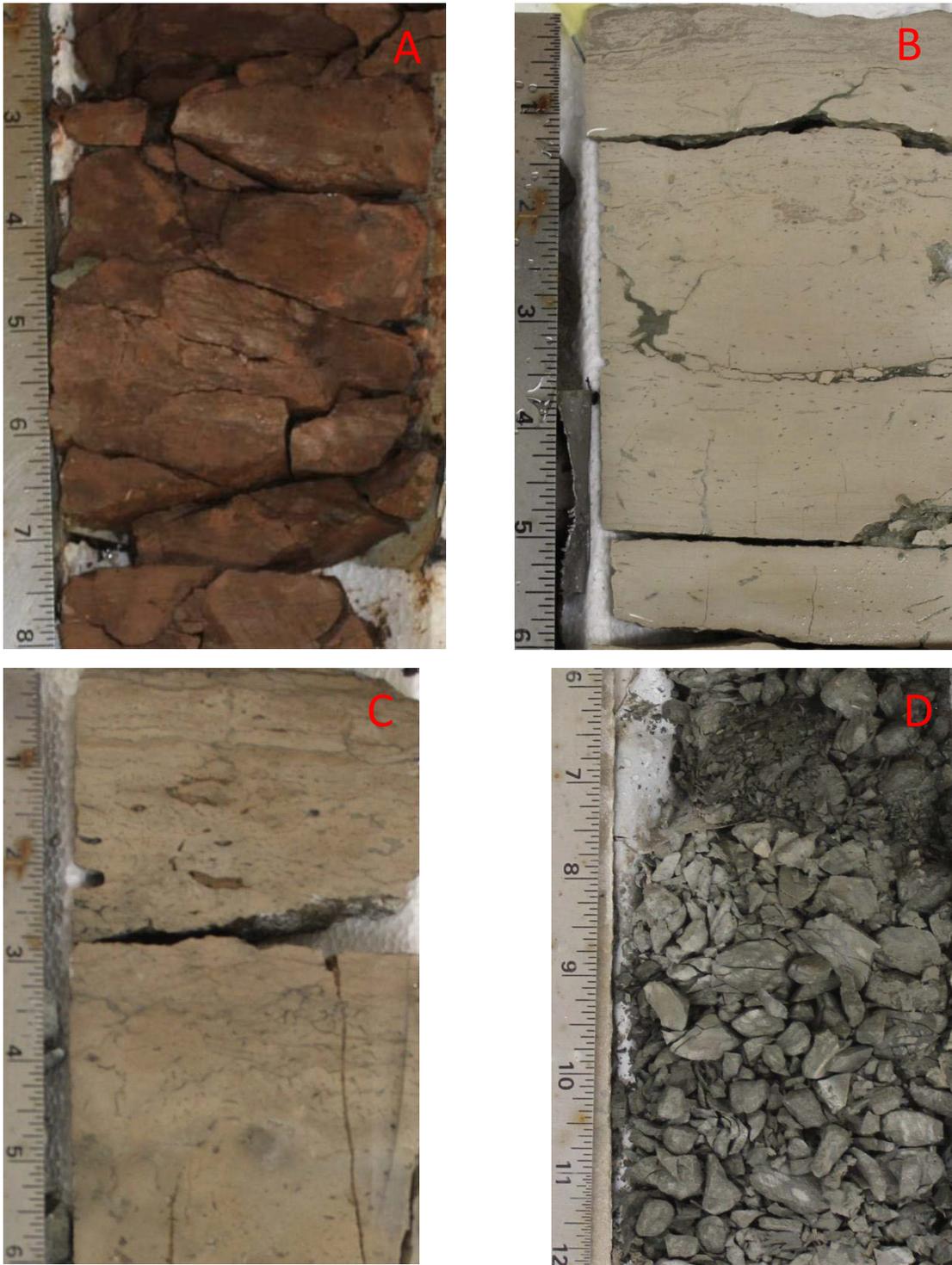


Figure 4 – Core photographs of the Dopita A #16; representative of the farthest points of regression in the LKC section: A) Red incipient paleosol indicative of a “full-cycle” regression to the point of subaerial exposure; Depth 3281’. B and C) Tan lime mudstone with birdseyes, *in situ* brecciation, and pore spaces filled with green calcitic shale indicative of a tidal-flat flat environment; Depths B-3258’6”, C-3324’. D) Green incipient paleosol with inferred poorly developed pedes indicative of a very shallow environment, with inferred marine plant life, though not subaerially exposed; 3294’.



Figure 5 - 'L' Zone - well sorted oolitic lime grainstone with a soilstone crust (red arrow); indicative of an unconformity. Approximate Depth – 3347'.



Figure 6 – ‘I’ Zone - carbonate clasts in a matrix of pale-green shale. Red arrow points to an accumulation of carbonate sediment below a larger carbonate clast; this indicates an abrupt sea-level fall resulting in water percolating through the rock, leaving sediment debris behind as it passes by larger clasts. Approximate depth – 3282’.



Figure 7 – ‘K’ Zone – lime grainstone with interparticle and vuggy porosity. Most the vugs (2-4mm in diameter) are filled with a moderately sorted, loosely consolidated and oil stained, carbonate sand. Unit contains lenses cemented by clear calcite and lack of readily visible porosity. Plug analysis results showed 5.5-6% porosity and permeability of .58-.64 mD; Approximate depth – 3326’.



Figure 8 – ‘I’ Zone - lime packstone to grainstone with low to high relief stylolites; interparticle, intraparticle, and moldic porosity (from partial dissolution of fusulinid and bryozoan fragments - from thin section analysis), sparse lenses with reduced porosity. Plug analysis showed 13.9% porosity and permeability of 3.93 mD; approximate depth – 3288’6”.



Figure 9 – ‘G’ Zone -well sorted oolitic sand shoal - grainstone, cross-stratified, with low-relief stylolites, indicative of a high-energy environment during a regression. This facies grades upward into a tidal flat environment in this cycle. Zone is completely oil saturated with 25% porosity and permeability of 2.68 md. Approximate depth – 3226’.



Figure 10 – 'I' Zone – lime biopackstone to grainstone with low to moderate relief stylolites with interparticle and moldic porosity (partially dissolved fusulinid and sparse bryozoan fragments – from thin section analysis). Most porosity is localized around stylolites. Plug analysis results showed 7.4% porosity and permeability of .71 mD. Approximate Depth – 3284'6".



Figure 11 – 'D' zone – lime biograinstone with low to moderate relief stylolites; contains interparticle, intraparticle, and vuggy porosity. Unit also contains fractures that were filled with calcite cement and partially dissolved. Plug analysis showed 9.1% porosity and permeability of 3.18 mD; Approximate depth – 3183'.



Figure 12 - 'A' Zone – lime biograinstone with an abundance of oncolites. Unit contains interparticle porosity and partial vuggy porosity. Two plugs were taken from the depths shown with the red arrows. Plug A results were 4.1% porosity with permeability of .0014 mD. Plug B results were 7.2% porosity and permeability of 4.23 m; Approximate depth – 3128'6".

Depositional Model

With the results compiled from all methods of study, a columnar section was constructed (Figure 13, or for large print see Plate 1) that shows the entire LKC in the core together with a brief lithologic description, presence of oil staining, porosity, fractures and inferred depositional environment. Also on the diagram are an inferred relative sea-level curve with complete cycles and minor cycles indicated.

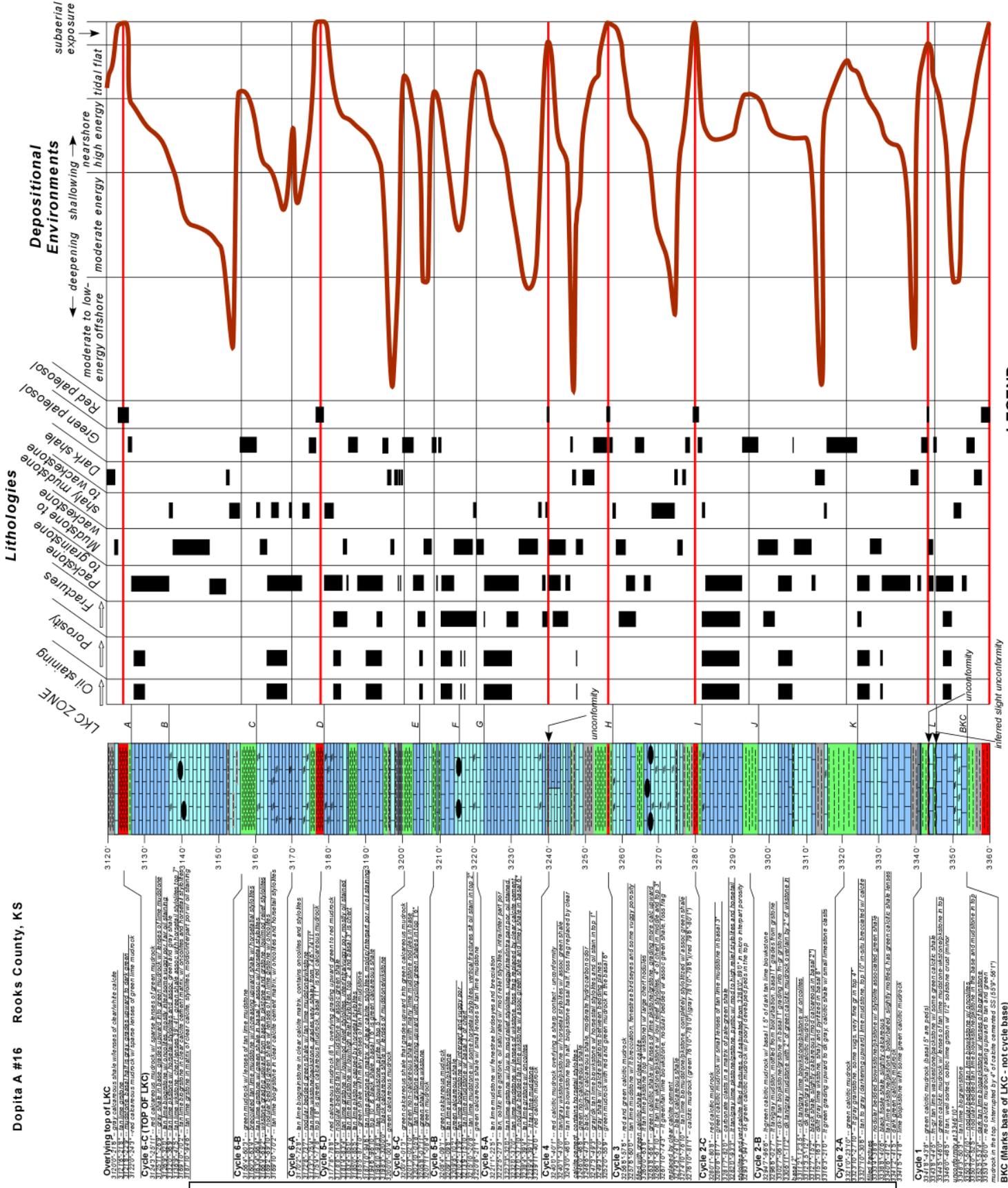


Figure 13 - Columnar section of the LKC groups based on the core analysis of the Dopita A #16 well in Rooks County, KS. - Larger Print is available, see Plate 2

Diagenetic History

Porosity development in carbonate rocks has been widely accepted as being secondary porosity primarily created or modified during the eogenetic stage attending meteoric subaerial exposure (Choquette and Pray, 1970). Partial enhancement or creation of secondary porosity can also occur through the mesogenetic (deep-burial) stage (e.g., Mazzullo and Harris, 1991). When carbonate sediments are deposited they commonly have primary porosities of 40-70% (Choquette and Pray, 1970). Post-depositional cementation reduces primary porosity, and burial-induced mechanical and chemical compaction also reduce porosity (Choquette and Pray, 1970). Analysis of the #16 core indicates that every zone in the LKC section was stylolitized during deep burial, and stylolitization generally increases rock density and likely reduces porosity even further (Mazzullo and Harris, 1991). The core analysis indicates that the only significant porosity present in the section is in rocks deposited in high-energy environments (packstones to grainstones) during regressions -- that is, in upward-shallowing highstand-systems tract deposits immediately beneath unconformities (see Figure 13). This observation suggests that the reservoirs were probably developed during the eogenetic stage attending periods of subaerial meteoric exposure. Porosity seemingly was the result of secondary dissolution of previously-precipitated, meteoric calcite cements. A similar mechanism of porosity development was suggested by (Watney et al., 1980) for LKC rocks elsewhere in Kansas.

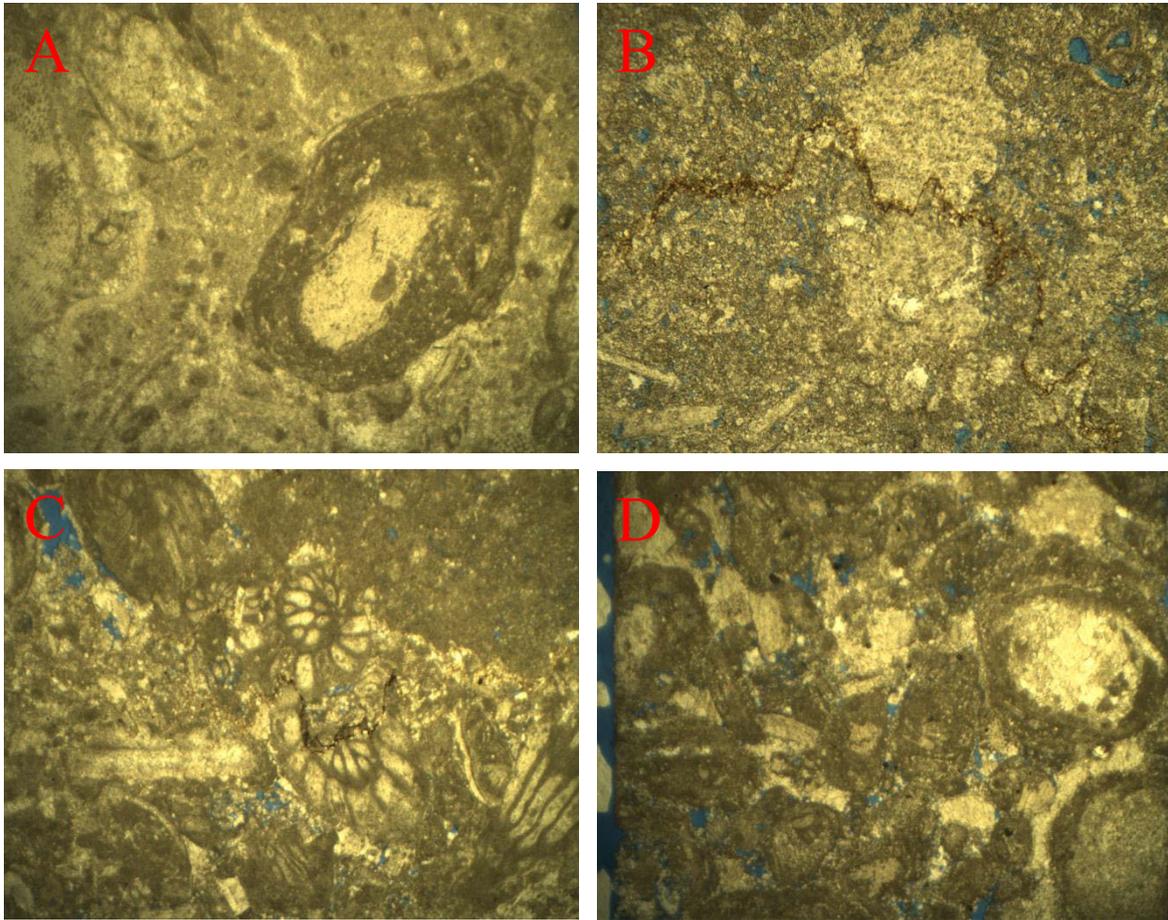


Figure 14 - Thin-section photomicrographs. **A)** 'I' Zone (plane light, 20X) - lime packstone to grainstone with no visible porosity. Rock contains many oncolites, fusulinids and some brachiopods and bryozoans. The primary porosity that was likely present has been filled with clear calcite cement that was not dissolved. Depth – 3292'0". **B)** 'I' Zone (plane light, 80X) - lime packstone to grainstone with mostly micro-interparticle porosity and some intraparticle and moldic (partly dissolved bryozoan fragments) porosity. Rock contains fusulinids and bryozoans and has dark oil staining along a moderate-relief stylolite. Depth – 3288'3". **C)** 'I' Zone (plane light, 20X) - lime packstone to grainstone with very sparse interparticle, intraparticle, and moldic (partial dissolution of fusulinid and bryozoan fragments) porosity. Rock contains sparse oil spotting and some dark oil staining concentrated along a low-relief stylolites. Primary porosity was filled with clear calcite cement that later was partly dissolved, producing the secondary porosity in the rock. Thin section contains forams, bryozoans, peloids, brachiopods, crinoids, and fusulinids. Depth – 3285'8". **D)** 'I' Zone (plane light, 20X) - lime packstone to grainstone with mostly interparticle and moldic porosity (partially dissolved fusulinid and sparse bryozoan fragments). Primary porosity was filled by clear calcite cement that was later partly dissolved, creating the secondary porosity in the rock. Depth – 3284'0".

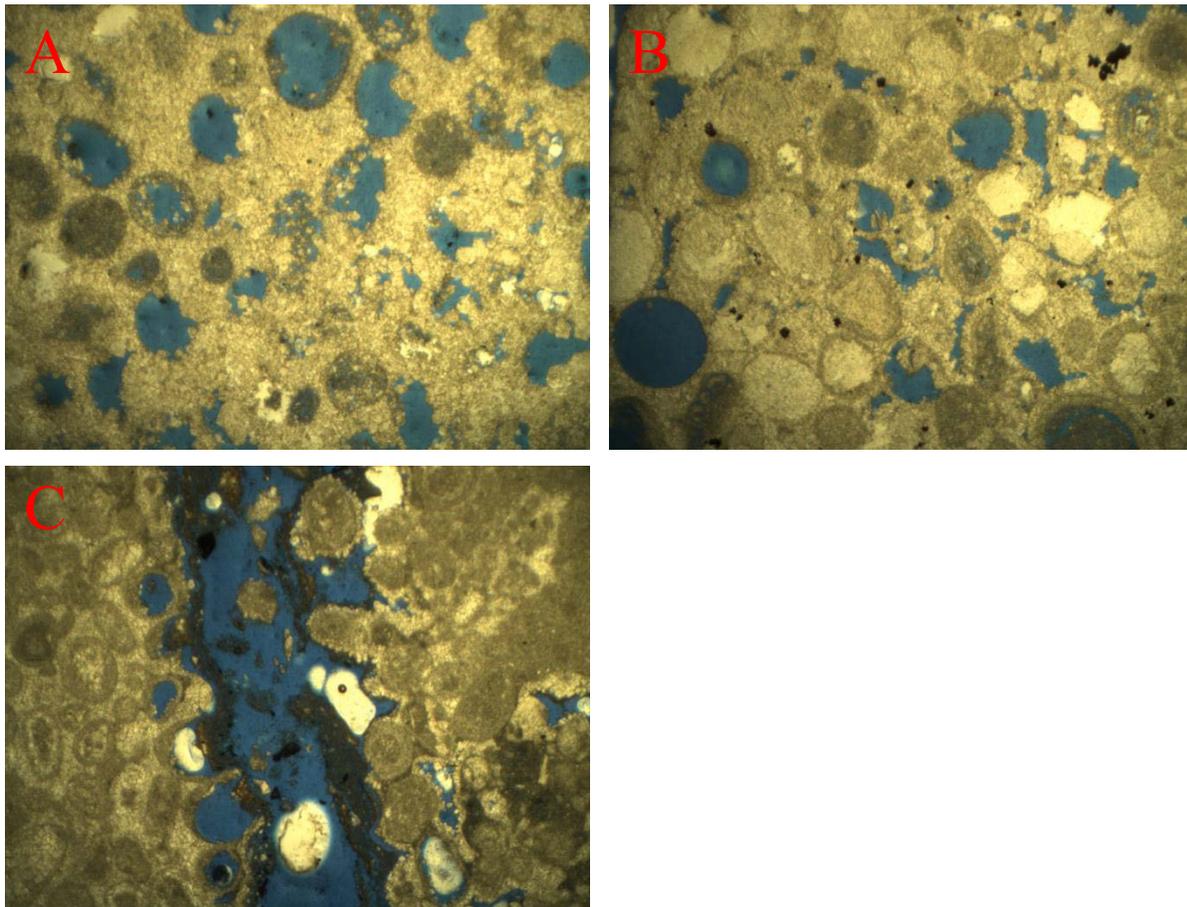


Figure 15 – Thin-section photomicrographs. **A)** ‘G’ Zone (plane light, 20X) - oolitic lime grainstone with oomoldic porosity and sparse interparticle porosity. Primary porosity was filled by clear calcite which was later dissolved, creating secondary porosity. Plug analysis showed 25. 2% porosity and permeability of 2.68 mD. Depth – 3226’4”. **B)** ‘G’ Zone (plane light,20X) - oolitic lime grainstone with both interparticle and oomoldic porosity. An abundance of oil spots are also present in the pores. Porosity appears to be approximately 50% interparticle and 50% oomoldic. Primary porosity was filled with clear calcite cement and later partially dissolved, creating secondary interparticle and oomoldic porosity. Depth – 3224’3”. **C)** ‘G’ Zone (plane light, 20X) - oolitic lime grainstone partially cemented by clear calcite. There is interparticle, oomoldic, and vuggy porosity. Most of the porosity present is moldic (oomolds) or vuggy. The primary porosity was filled with clear calcite cement and later dissolved, creating secondary porosity. Fractures are visible throughout the thin section and oil is present in some pores. Plug analysis showed 10.6% porosity and permeability of 111 mD (likely a pore space connected the length of the plug and misrepresentative of the overall formation). Depth – 3222’9”.

Porosity and Permeability Analysis of Plugs from Prominent Dopita Field Reservoirs



SUMMARY OF ROUTINE CORE ANALYSES RESULTS

Vacuum Dried at 180°F Net Confining Stress: 800 psi

TREK AEC
Dopita A#16 Well

Kansas

LKC Zone	Sample Depth, feet	Permeability, millidarcys		Porosity, percent		Grain Density, gm/cc
		to Air	Klinkenberg	Ambient	NCS	
A	3128'2"	0.0014	0.0003	4.1	4.0	2.71
A	3128'9"	4.23	3.20	7.2	7.1	2.71
D	3181'6"	3.63	2.92	8.3	8.2	2.71
D (v)	3182'7"	3.18	2.56	9.1	9.1	2.70
D (v)	3193'9"	0.013	0.0049	0.5	0.5	2.71
E (v)	3211'2"	10.4	8.40	13.6	13.5	2.74
Lower E (v)	3212'10"	0.0011	0.0002	3.3	3.2	2.71
G (v)	3223'3"	111.	99.2	10.6	10.5	2.71
G (v)	3226'0"	2.68	2.16	25.2	25.1	2.71
I	3284'2"	0.713	0.531	7.4	7.3	2.72
I	3288'4"	3.93	3.17	13.9	13.8	2.71
J	3303'10"	0.024	0.011	9.5	9.5	2.73
J (v)	3304'6"	0.113	0.067	11.8	11.7	2.73
K	3325'5"	0.581	0.429	5.5	5.4	2.71
K (v)	3326'4"	0.643	0.480	6.0	5.9	2.70
L	3348'10"	0.210	0.136	2.6	2.6	2.70
	Average values:	8.83	7.70	8.7	8.6	2.71

Figure 16 – Routine core analysis of porosity and permeability from permeability plugs
(Weatherford Laboratories, 2013).

The plugs used for the porosity and permeability analysis (Figure 16) all came from high-energy packstones and grainstones within the labeled zones and were chosen by the presence of porosity and/or oil staining when analyzing the core. Within each zone, both porosity and permeability decrease upward, with the majority of the porosity and permeability concentrated in the middle of the zone. This is noticeable in Figure 16. It would appear that the 'E' zone is in an exception in the chart, though the two plugs were extracted from two separate, incomplete, sub-cycles (Figure 13 or Plate 1).

Discussion and Conclusions

The study showed that reservoir development in the Dopita A #16 well is present only in high-energy carbonate rocks in the cycles (highstand-systems tract packstones to grainstones) deposited during regression. Secondary porosity formation in these rocks resulted from meteoric dissolution beneath unconformities in the section. The high frequency glacial eustasy inferred during this time would give each zone in the LKC section the potential for developing porosity. Examination of the columnar section (Figure 13 and Plate 1) indicates that porosity development is dependent on regression. The extent of the regression appears to not have a direct correlation to the amount of porosity. As long as sea-level regressed to a tidal flat environment or shallower, porosity was able to develop by meteoric dissolution beneath unconformities.

It has been estimated that an oil-producing Lansing-Kansas City reservoir in Kansas with three or more wells would have a 75% chance of having some economic potential if waterflooded (Koudele and Wilhite, 1990). In the north end of the Dopita Field a waterflood has already been implemented and is proving to be very successful. Current estimates point to a

secondary recovery of .82:1 ratio to their primary recovery (DeGood, 2013). It is highly likely that a flood will work in the south end of the field. Accidental casing leaks in the south end of the field have doubled the monthly production of some leases and immediately dropped production rates when the leaks were fixed.

Plans have already been made to make the Dopita A #16 an injection well as part of a waterflood proposal. Other wells in the field will be converted to injection wells and more wells will possibly be drilled.

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